RESEARCH ARTICLE ENVIRONMENTAL STUDIES

Quantifying methane emissions from the largest oil-producing basin in the United States from space

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Abstract

Using new satellite observations and atmospheric inverse modeling, we report methane emissions from the Permian Basin, which is among the world's most prolific oil-producing regions and accounts for >30% of total U.S. oil production. Based on satellite measurements from May 2018 to March 2019, Permian methane emisi≡ sions from oil and natural gas production are estimated to be ① 7 ± 0.5 Tg a⁻¹, representing the largest methane flux ever reported from a U.S. oil/gas-producing region and are more than two times higher than bottom-up inventory-based estimates. This magnitude of emissions is 3.7% of the gross gas extracted in the Permian, i.e., ~60% higher than the national average leakage rate. The high thane leakage rate is likely contributed by extensive venting and flaring, resulting from insufficient infrastructure to process and transport natural gas. This work demonstrates a high-resolution satellite data-based atmospheric inversion framework, providing a robust top-down analytical tool for quantifying and evaluating bregional methane emissions.

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INTRODUCTION

Methane is a potent greenhouse gas with a relatively short average atmospheric residence time of about a decade and is also a precursor of tropospheric ozone (1). The emission-based radiative forcing for methane (including effects on tropospheric ozone and stratospheric water vapor) is 0.97 W m⁻² since preindustrial times, which is about 60% of that for CO₂ (2). Roughly a third of the contemporary anthropogenic methane emissions come from the fossil fuel energy sector worldwide (oil, natural gas, and coal) (~100 to 180 Tg a⁻¹) (3, 4, 5). Curbing anthropogenic methane emissions, including those from the oil/gas sector, is considered an effective strategy to slow the rate of near-term climate warming (1). However, the rapid increase in oil and natural gas (O/G) production in the United States since around 2005, driven primarily by hydraulic fracturing and horizontal drilling, has led to major concerns about increasing methane emissions and adverse climate impacts (6). By upscaling data collected from field measurements in some of the largest O/G production basins in the United States, Alvarez *et al.* (7) estimated 13 Tg

annual methane emissions from the national O/G supply chain for 2015, which is 60% higher than the official estimates by the U.S. Environmental Protection Agency (EPA) ($\underline{8}$). The largest discrepancy was found in the O/G production segment where the estimate by Alvarez *et al.* ($\underline{7}$) (7.6 Tg a⁻¹) was more than two times that by EPA, which relies on inventory-based estimates (3.5 Tg a⁻¹) ($\underline{8}$).

While field measurements provide in-depth information about a particular site or area, it is often challenging to expand the measurement capacity to observe a diverse set of targets distributed globally over longer periods of time. Additional challenges exist for areas that are difficult to access for technical or proprietary reasons. On the other hand, global satellite observations of column atmospheric methane offer a unique vantage point to identify emission hot spots and quantify regional emissions (2). Using data from SCanning Imaging Absorption spectroMeter for Atmospheric CHartographY (SCIAMACHY) satellite observations averaged between 2003 and 2009, Kort *et al.* (10) found large anomalous methane levels from the Four Corners region in the United States, with total methane emissions associated with natural gas, coal, and coalbed sources estimated as 0.59 ± 0.08 Tg a⁻¹. While the

SCIAMACHY data were fairly limited in spatial resolution (30 km × 60 km) and measurement precision [30 parts per billion in volume or (ppbv)] ($\frac{9}{2}$), it was the first time that satellite observations were used to quantify a dense O/G-related methane emission hot spot. This finding also led to several dedicated airborne studies to better understand methane sources in the region ($\frac{11}{11}$, $\frac{12}{12}$), which reported methane fluxes comparable to the satellite-based estimate ($\frac{10}{10}$).

Here, we demonstrate and exploit the capability of a recent space-borne sensor, the Tropospheric Monitoring Instrument (TROPOMI), to map atmospheric methane enhancements in the United States and quantify emissions from the Permian Basin (Fig. 1), which has become one of the world's most prolific oil-producing regions in recent years due to advances in drilling technologies. Located in New Mexico and Texas in a region of ~400 km × 400 km, Permian is currently the largest oil-producing basin in the United States. In 2018, the Permian Basin produced 5.5 × 10⁵ m³ (or 3.5 million barrels) of crude oil and 3.2×10^8 m³ (or 11 billion feet³) of natural gas every day (~30 and ~10% of the U.S. national totals, respectively), which was 4 and 2.5 times their corresponding levels in 2007 (around the time of SCIAMACHY observations) (Fig. 2) (13). While the surging production in the Permian Basin and its importance in the U.S. oil boom during the last decade have been widely covered in mass media (14), the scale of associated methane emissions from this critical O/G basin is unknown, despite reports of increased flaring and venting activity (15).

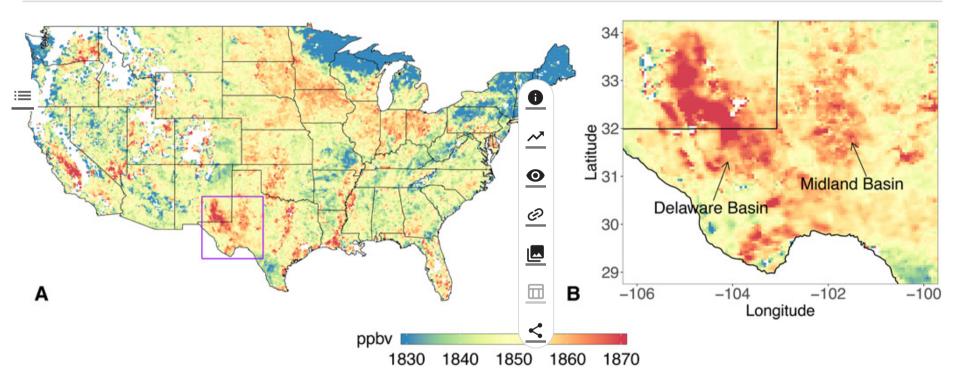
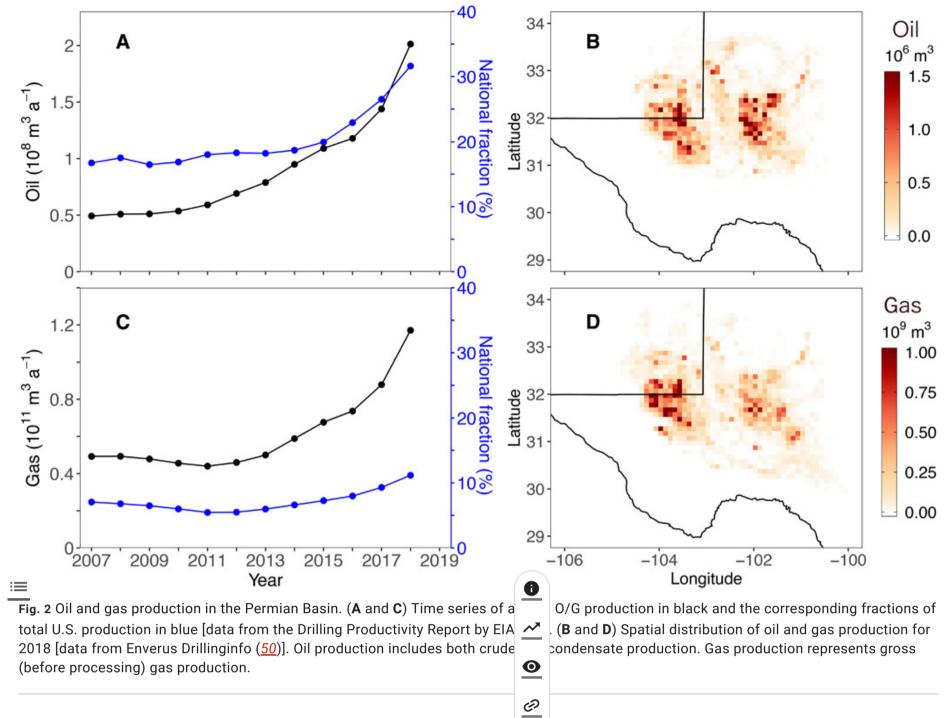


Fig. 1 Satellite observations of the Permian methane anomaly. TROPOMI satellite data derived elevation-corrected column methane mixing ratio for (**A**) the conterminous United States and (**B**) the Permian Basin containing the Delaware and Midland sub-basins. White shading represents missing data. Purple boundary in (A) indicates the study domain encompassing the Permian Basin. Methane averages are computed from monthly means of TROPOMI measurements during May 2018 and March 2019.



Using 11 months of recent data acquired by TROPOMI during centration anomaly over the Permian Basin and quantify the a 🛄 iated methane emissions with a state-of-the-art atmospheric inverse modeling framework. TROPOMI was laun 🔟 l in October 2017 onboard the European Space Agency's Sentinel-5P satellite and provides column atmosphel 🤜 lethane measurements with higher spatial reso-

-2019, we focus on the distinct methane con-

lution (7 km \times 7 km at nadir) and precision (0.6%) than was previously available (<u>16</u>), providing near-daily global coverage with its large 2600-km-wide swath (<u>17</u>). Our integrated satellite-based approach provides new insights into the dynamic landscape of O/G-related methane emissions in the United States and should pave the way forward toward routine quantification, monitoring, and evaluation of methane emissions from source regions distributed globally.

RESULTS

Satellite observations of the Permian methane anomaly

Figure 1A shows a map of column-averaged dry-air methane mixing ratio over the conterminous United States, retrieved from TROPOMI measurements, with correction for the topography effect (denoted as XCH_4^t ; see Materials and Methods). The data are averaged from May 2018 to March 2019. Substantial enhancements of XCH^t₄ relative

to the surrounding background, up to ~30 ppbv, are found over the Permian Basin, indicating strong methane emissions. Other notable enhancements are observed in California's central valley, coastal Southeast, and the Mississippi River Valley, likely associated with anthropogenic (agriculture, dairy) and natural (wetland) sources. The elevated methane levels in central California were also seen earlier in the SCIAMACHY analysis (<u>10</u>).

The methane enhancements over the Permian Basin show a characteristic two-branch pattern, which aligns with the two major O/G production sub-basins, the Delaware basin to the west and the Midland basin to the east (Fig. <u>1B</u>). The enhancement over the Delaware basin, where extensive new exploitation has taken place during the last 5 years (<u>18</u>) (fig. S1), is larger than that over the Midland basin (<u>Fig. 1B</u>). Intensive O/G production activity in these two sub-basins is also captured by satellite observations of radiant heat from gas flaring [Fig. 3A; nighttime observations by the Visible Infrared Imaging Radiometer Suite (VIIRS)] and NO₂ tropospheric column densities (Fig. 3B; daytime observations by TROPOMI). Flaring is a common practice in O/G operations to burn off unwanted or excess gas, and NO₂ is a gaseous pollutant released during gas flaring and other combustion activities in O/G fields

(<u>19</u>, <u>20</u>). On the basis of measurements by the VIIRS instrument onboard the Suomi National Polar-orbiting Partnership satellite, we estimate an average flaring rate of 5.9 ± 1.2 billion m³ a⁻¹ during the period of this study, about 4.6% of the gross gas production (see text S1). A fourfold increase in flaring intensity since 2012, observed by the VIIRS instrument, is indicative of the rapid growth in O/G production across the Permian Basin (fig. S1).

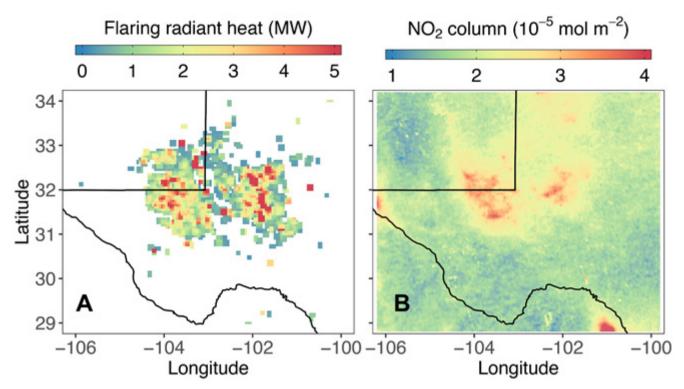


Fig. 3 Satellite observations of gas flaring radiant heat and NO₂ tropospheric column density over the Permian Basin. (A) Gas flaring radiant heat is the annual average of 2018 measured by the VIIRS satellite instrument, and (B) NO₂ tropospheric column density is the 3-month aver image of 2018) measured by the TROPOMI instrument, image of 2018) measured by the TROPOMI instrument, image of 2018 measured by the TROPOMI instrument in the trop of 2018 measured by the TROPOMI instrument, image of 2018 measured by the TROPOMI instrument in the trop of 2018 measured by the trop of 2018 measu

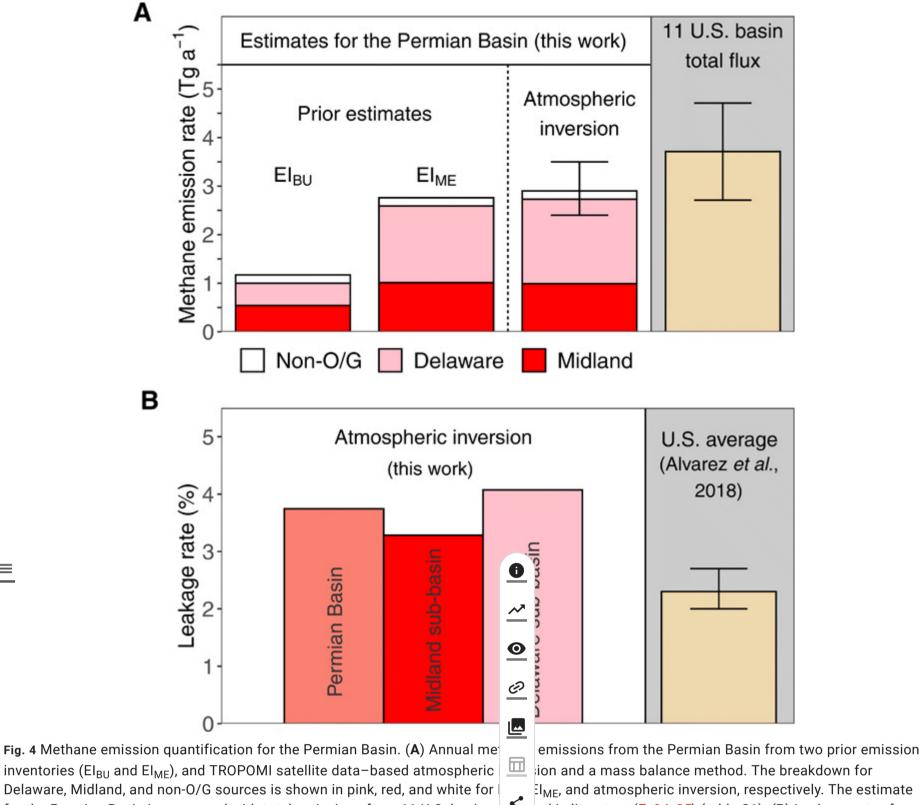
Methane emission quantification

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We quantify the methane emission rate from the Permian Basi \checkmark d its spatial distribution with atmospheric inverse modeling, which optimizes spatially resolved methane e TROPOMI observations and the prior emission estimate follow mize monthly methane emission rates resolved at $0.25^{\circ} \times 0.31$ in rates by drawing information from the Bayesian rule. The inversion seeks to optimize the Permian Basin and the surrounding region $(29^{\circ}-34^{\circ}N)$, \checkmark $-106^{\circ}W$). The solution to the optimization is found analytically with closed-form characterization of the error statistics (3). An atmospheric transport model (a nested version of GEOS-Chem over North America with a $0.25^{\circ} \times 0.3125^{\circ}$ horizontal resolution) (21) is used as the forward model to relate atmospheric methane columns with ground-level emissions in the study domain and the contributions from outside the domain. The optimization by the inversion significantly reduces the observation-model mismatch with decreased root mean square error (prior, 23 ppbv; posterior, 14 ppbv) and increased correlation (R; prior, 0.30; posterior, 0.62) (fig. S2). See Materials and Methods for more details about the configurations of the inverse modeling including error accounting and prior information.

When aggregating monthly spatially resolved posterior emissions to the basin-level annual average, we find a methane emission flux of 2.9 ± 0.5 Tg a⁻¹ from the Permian Basin ($30^{\circ}-34^{\circ}N$, $101^{\circ}-105^{\circ}W$) (Fig. 4A; see Materials and Methods for the uncertainty analysis). This estimate is more than a factor of 2 larger than the bottom-up estimate based on an extrapolation of EPA greenhouse gas inventory data (EI_{BU}, 1.2 Tg a⁻¹; see Materials and Methods) (Fig. 4A), suggesting that current methane emissions in the Permian are underrepresented in national bottom-up emission inventories (22). Our inversion result is in close agreement with a basin-level estimate based on extrapolation of limited ground-based site-level measurements in the Permian (EI_{ME}, 2.8 Tg a⁻¹) (Fig. 4A). It should be noted that these site-level measurements were primarily conducted in the New Mexico portion of the Permian Basin and covered only a small fraction of production sites (see Materials and Methods and text S2). As a comparison, we also apply a fast mass balance method following Buchwitz *et al.* (23) to estimate basin-level emissions, which yields an annual mean emission rate of 3.2 ± 2.0 Tg a⁻¹ for the Permian Basin. This result is consistent with that derived from a full atmospheric inversion. Despite the large uncertainty of the mass balance method, this data-driven approach provides an independent estimate of emissions derived primarily using TROPOMI data (see text S3 for more discussion).



< for the Permian Basin is compared with total emissions from 11 U.S. basins k ted in literature (7, 24, 25) (table S1). (B) Leakage rates for the Permian Basin and two sub-basins, in comparison with the average leakage reported for the entire United States (Z).

Removing the non-O/G sources (0.2 Tg a^{-1}) from the total flux obtained via the inversion (2.9 Tg a^{-1}), we estimate the methane emissions related to O/G activity to be 2.7 Tg a^{-1} in the Permian Basin. Put in the context of national emissions, this value is approximately one quarter of total emissions from all U.S. oil and gas production areas in 2015 (10.9 Tg a⁻¹, including emissions from production, gathering, and processing, which largely occur in the production areas) (Z). Our estimated emission rate for the Permian is significantly higher than those reported in the literature for other major U.S. O/G-producing basins. Table S1 summarizes methane emission estimates for 11 U.S. basins (7, 24, 25) from previous aircraft-based studies [i.e., Haynesville (24, 26), Barnett (24, 27), Northeast Pennsylvania (<u>26</u>, <u>28</u>), Southwest Pennsylvania (<u>25</u>), San Juan (<u>12</u>), Fayetteville (<u>26</u>, <u>29</u>), Bakken (<u>24</u>, <u>30</u>), Uinta (<u>31</u>), Weld (<u>32</u>), West Arkoma (<u>26</u>), Eagle Ford (<u>24</u>), and the Denver Basin (<u>24</u>)]. Our estimate for the Permian (2.7 Tg a^{-1}) is about a factor of 4 higher than the largest methane emissions from these previously reported O/G basins [i.e.,

Eagle Ford, 0.73 Tg a^{-1} (<u>24</u>)] and is even comparable to the 11-basin sum (3.7 Tg a^{-1}) (<u>Fig. 4A</u> and table S1). This comparison with recent literature indicates that the Permian Basin is likely the largest observed methane-emitting O/G basin in the United States and a substantial contributor to national O/G-related emissions.

Distribution of methane emissions

High-resolution observations from TROPOMI enable us to resolve methane emissions at an unprecedented spatial and temporal resolution, relative to the previous generation of satellite instruments such as the Greenhouse gases Observing SATellite (GOSAT) and SCIAMACHY (9). Figure 5 presents the spatial distribution of methane emissions in the Permian Basin at about a quarter-degree resolution derived from our atmospheric inversion. Compared to the prior inventory EI_{BU}, our inversion finds larger methane emissions near the center of the Delaware and Midland sub-basins. Sensitivity inversions further show that this spatial pattern is robust against prior emissions

of varied magnitudes and distributions (fig. S3), demonstrating that it is primarily informed by satellite observations.

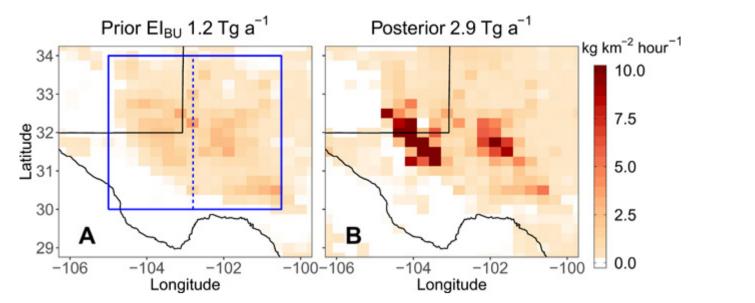


Fig. 5 Spatial distribution of methane emission rates in the Permian Basin. (**A**) Bottom-up emission inventory El_{BU} extrapolated from EPA greenhouse gas inventory data (prior). (**B**) TROPOMI observation-derived emissions using Bayesian atmospheric inverse modeling (posterior). The prior and posterior basin-total emissions, indicated on top of the figure, are computed over the area enclosed by the solid blue boundary, with contributions from two sub-basins, the Delaware (left of the dashed line) and Midland (right of the dashed line).

The spatial distribution of methane emissions derived from inversion is closely correlated with that of gross gas production (R = 0.78), but to a lesser degree with that of oil production (R = 0.53) and that of the well number denity (R = 0.31) (fig. S4). Similarly, when we sum up the O/G-relation emissions between Delaware and Midland (1.7/1.0 Tg a⁻¹ = 1) ane emissions between Delaware and Midland (1.7/1.0 Tg a⁻¹ = 1) compared to that of oil production (1.0) and well number density suggests that with these wells (e.g., gathering stations), which have been de ters in the Permian Basin.

In addition to the spatial distribution, our monthly inversion a provides information about the temporal variation of methane emissions during the 11 months of observation of g. S5). Although the inversion's ability to resolve the spatial distribution of emissions varies from month to month because of uneven monthly sampling of TROPOMI (fig. S5), our inversion ensemble (table S2 and fig. S5) generally results in consistent monthly basinlevel emission estimates (see also uncertainty analysis in Materials and Methods). We speculate that high emissions in December 2018 may be related to a very low in-basin gas price toward the end of 2018, resulting from insufficient gas gathering and transmission capacity in the Permian Basin (<u>33,34</u>). That said, we do not find an apparent increasing trend in methane emissions, although natural gas production from the Permian Basin increased steadily by ~20% during the overlapping 11-month period (fig. S6). Further investigation is required to delineate factors controlling the temporal variations of O/G-related methane emissions.

DISCUSSION

Using an inverse analysis of TROPOMI satellite observations, we estimate a total methane flux of 2.9 ± 0.5 Tg a⁻¹ in

the Permian Basin, with 2.7 Tg a⁻¹ coming from O/G-related activity. Methane losses of this magnitude represent a waste of an important resource; for instance, this is enough natural gas to supply 7 million households in the state of Texas (<u>35</u>). Moreover, the 2.7 Tg a⁻¹ methane emitted in Permian results in the same radiative forcing as ~260 Tg a⁻¹ CO₂ over a 20-year time horizon (86 Tg CO₂ a⁻¹ over a 100-year time horizon) (global warming potential of 96 for 20 years and 32 for 100 years) (<u>7</u>, <u>36</u>), about the same as annual CO₂ emissions from the entire U.S. residential sector (290 Tg CO₂ a⁻¹ in 2017) (<u>22</u>).

Our estimate (2.7 Tg a⁻¹) equates to a production-normalized (73 Tg CH₄ a⁻¹, derived from 127 billion m³ a⁻¹ natural gas production during the study period using 80% methane content by volume) emission rate (or methane leakage rate) of $3.7 \pm 0.7\%$, which is ~60% higher than the national average of $2.3 \pm 0.3\%$ ($\underline{7}$) (Fig. 4B). The leakage rate is even higher for the rapidly developing Delaware sub-basin (4.1%). Comparable high leakage rates have also been reported in other oil production–focused basins such as the Bakken ($\underline{24}$) (table S1), but these basins produce much lower natural gas than the Permian Basin does. Previous studies summarized in table S1 show an inverse relation-

ship between the basin-level leakage rate and gas production (<u>24</u>); however, the Permian Basin is an outlier with high oil production, high gas production, and a high leakage rate.

Overall, the high leakage rate in the Permian Basin appears to be associated with insufficient infrastructure for natural gas gathering, processing, and transportation (34, 37), leading to extensive venting and flaring (Fig. 3), which contributes to high methane emissions. The greater profitability of oil production contributes to a lack of investment in natural gas takeaway capacity, which, in turn, has resulted in excessive supply of associated gas and a very low in-basin gas price in the Permian (34). In addition, with the rescinding of U.S. federal requirements on gas capture and fugitive emissions in 2018, current regulations on O/G methane emissions in the Permian Basin are less stringent at both federal and state levels (see text S4). All these factors may increase the incentive for operators to vent and flare their product. On the other hand, the higher-than-average leakage rate in the Permian Basin implies an opportunity to reduce methane emissions in this rapidly growing oil and gas–producing region, through better design, effective management, regulation, and infrastructure development.

MATERIALS AND METHODS

TROPOMI methane observations

We use daily column-averaged dry air column methane mixing ratio (XCH₄) data retrieved from TROPOMI measurements (<u>38</u>) between May 2018 and March 2019. TROPOMI, onboard the polar-orbiting Sentinel-5 Precursor satellite, is a push-broom imaging spectrometer that provides near-daily global coverage with a swath width of $^{\circ}$ 500 km and a nadir ground pixel size of 7 km × 7 km at approximately 13:30 local overpass time (<u>17</u>). The retrieval taneously inferring methane concentrations and physical scattering properties, using the oxygen A-band in the band in the short-wave infrared (SWIR) (<u>39</u>). Only high-qualitic conditions are used in this study (as indicated by the retrieval These measurements are filtered for solar zenith angle (<70°), (1 SD of surface elevation <80 m within 5-km radius), and low (<u>40</u>).

The TROPOMI XCH₄ product is further corrected for any know \checkmark rieval biases (40). The errors in the TROPOMI XCH₄ measurements have been assessed against GOSAT XCH₄ data (38) and were found to correlate with surface albedo. A global bias correction linearly dependent on surface albedo was then derived and applied to the TROPOMI data (40). This bias-corrected TROPOMI XCH₄ product is used in this study. Negligible correlation of errors with other retrieved parameters (e.g., aerosol optical thickness) was found in the assessment. Validation with independent ground-based measurements from the Total Column Carbon Observing Network shows that the bias-corrected TROPOMI XCH₄ has a bias of -4.3 ± 7.4 ppbv, improved upon the uncorrected XCH₄ and other retrieved parameters for the subset of TROPOMI data over the domain of this study. We find no correlation with albedo ($R^2 = 0.00$) and a negligible correlation with aerosol optical thickness ($R^2 = 0.07$), supporting the idea that the XCH₄ enhancement over the Permian Basin (Fig. 1B) is robust.

Figure S7A shows the average XCH₄ over the conterminous United States and the Permian Basin between May 2018

and March 2019 before the topographical correction. We derive the elevation-corrected methane column (XCH_4^t) shown in Fig. 1 by applying a third-order polynomial correction fitted over the U.S. domain following Kort *et al.* (*10*). The mass balance method uses the elevation-corrected data (XCH_4^t) for emission quantification, while the inversion method uses XCH₄ (bias-corrected) directly obtained from the data product, because the topography effect is taken care of by the atmospheric transport model.

Atmospheric inverse modeling

We perform an inverse analysis of TROPOMI observations to derive optimized estimation of monthly methane emissions at $0.25^{\circ} \times 0.3125^{\circ}$ horizontal resolution in the Permian Basin. Quantification of emissions at this combination of relatively high spatial and temporal resolution, not achievable with previous generations of satellite observations such as from GOSAT or SCIAMACHY, is enabled by higher-resolution TROPOMI satellite observations (<u>41</u>). Figure S7B shows that the Permian Basin is well sampled by TROPOMI during the study period, likely because of frequent cloud-free conditions in the region. A total of ~200,000 TROPOMI XCH₄ retrievals within the study domain (29°–34°N, 100°–106°W) between May 2018 and March 2019 are used for the inversion.

Let \mathbf{x} be the state vector that we seek to optimize through inversion, including a gridded ensemble of methane emissions and an additional element representing the regional model bias in XCH₄. The regional model bias term (a monthly scalar uniform over the inversion domain) is necessary to account for spatially uniform biases caused by imperfect lateral boundary condition and emission errors outside the study domain. The inversion solves for an optimal estimate of \mathbf{x} by minimizing the following cost function

$$J(\boldsymbol{x}) = (\boldsymbol{x} - \boldsymbol{x}_{\mathrm{A}})^{T} \mathbf{S}_{\mathrm{A}}^{-1} (\boldsymbol{x} - \boldsymbol{x}_{A}) + (\boldsymbol{y} - \mathbf{K}\boldsymbol{x})^{T} \mathbf{S}_{\mathrm{O}}^{-1} (\boldsymbol{y} - \mathbf{K}\boldsymbol{x})$$
(1)

where TROPOMI XCH₄ observations are assembled in y, x_A is the prior estimate of x, S_A is the prior error covariance matrix, S_O is the observational error covariance matrix, and K is the Jacobian matrix describing the sensitivity of XCH₄ to emissions and the regional model bias ($\partial y / \partial x$).

Minimization of Eq. 1 at $\nabla_{\mathbf{x}} J(\mathbf{x}) = 0$ yields the posterior estimation $(\widehat{\mathbf{x}})$, the posterior error covariance matrix $(\widehat{\mathbf{S}})$, and the averaging kernel matrix (A) (<u>42</u>)

$$\widehat{\boldsymbol{x}} = \boldsymbol{x}_{\mathrm{A}} + \boldsymbol{S}_{\mathrm{A}} \boldsymbol{\mathrm{K}}^{T} (\boldsymbol{\mathrm{K}} \boldsymbol{\mathrm{S}}_{\mathrm{A}} \boldsymbol{\mathrm{K}}^{T} + \boldsymbol{\mathrm{S}}_{\mathrm{O}})^{-1} (\boldsymbol{y} - \boldsymbol{\mathrm{K}} \boldsymbol{x}_{\mathrm{A}})$$
(2)

$$\widehat{\mathbf{S}} = \left(\mathbf{K}^T \mathbf{S}_{\mathrm{O}}^{-1} \mathbf{K} + \mathbf{S}_{\mathrm{A}}^{-1}\right)^{-1}$$
(3)

$$\mathbf{A} = \mathbf{I}_n - \widehat{\mathbf{S}} \mathbf{S}_{\mathbf{A}}^{-1} \quad \textcircled{\bullet} \tag{4}$$

Here, \mathbf{I}_n is an identity matrix where *n* is the dimension of the s grees of freedom for signal (DOFS), quantifies the number of p state vector. \mathbf{v} . The trace of **A**, often called as the des of information constraining the *n*-dimensional \mathbf{v}

To solve for Eqs. 2 to 4, the prior estimate (\mathbf{x}_A) for gridded met of information, we create two gridded emission inventories for tion (EI_{BU}) and the other based on extrapolation of ground-bases scriptions of the inventories). Both emission inventories are the set inversion, while we use EI_{ME} in a sensitivity inversion to evaluate the impact of the prior estimate (PI_EI_{ME}; see table S2). We perform further evaluations using prior emissions constructed by disaggregating the total O/G-related emission flux from EI_{BU} with varied spatial proxies (i.e., well count, PI_EI_{well}, natural gas production, PI_EI_{gas}, and oil production, PI_EI_{oil}) (table S2 and fig. S3).

The difference between the EI_{BU} and EI_{ME} (Fig. 5A and fig. S3A) measures the uncertainty of our prior knowledge, and we thus specify prior errors (S_A) for emissions as the absolute difference between EI_{BU} and EI_{ME} . We also specify the prior error for the regional model XCH₄ bias as 10 ppbv. To test the sensitivity to prior errors, we perturb S_A in two sensitivity inversions by doubling (PE × 2) or halving (PE × 0.5) prior errors (table S2). S_O is constructed with the residual error method (43), which results in an error averaged at ~11 ppbv. Both S_O and S_A are taken to be diagonal matrices. We also perform a sensitivity inversion to test the impact of error correlations with off-diagonal terms specified following Cusworth *et al.* (44) (OE_Cor; see table S2).

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A nested version of the GEOS-Chem chemical transport model (12.1.0) is used as the forward model in the inversion to link XCH₄ to surface emissions. To account for the vertical sensitivity of the satellite instrument, we compute simulated XCH₄ by applying TROPOMI averaging kernels to simulated methane vertical profiles. We construct the Jacobian matrix **K**, column by column, with simulations perturbing each state vector element independently. The simulations are performed over North America and adjacent oceans driven by GEOS-FP–assimilated meteorological data from the NASA Global Modeling and Assimilation Office on a $0.25^{\circ} \times 0.3125^{\circ}$ horizontal grid and 47 vertical layers (~30 layers in the troposphere) (*21*). The boundary conditions for the nested-grid simulation are from a $4^{\circ} \times 5^{\circ}$ global simulation from May 2018 to March 2019 driven by GEOS-FP meteorological fields. Note that methane emissions and sinks used in this simulation are optimized with previous-year (2010–2017) GOSAT satellite data following Maasakkers *et al.* (*3*). Such generated boundary conditions may be biased (i.e., unable to capture the growth of global methane concentrations; see fig. S9), and we account for it by introducing a monthly regional

model bias term in the inversion. The retrieved regional model biases may vary with the extent of the inversion domain. To test this sensitivity, we also perform an inversion with a larger spatial domain $(27^{\circ}-36^{\circ}N, 98^{\circ}-108^{\circ}W)$ (Bg_Large; see table S2).

Inversion uncertainty

The posterior error covariance matrix (\hat{S} , Eq. 2) and averaging kernel matrix (A, Eq. 3) evaluate the uncertainty of an inversion solution given inversion parameters (e.g., S_A , S_O , forward model). Figure S5 shows monthly posterior errors for basin-level emissions (derived from $\widehat{\mathbf{S}}$) and corresponding DOFS (trace of A) from our base inversion. Overall, the posterior errors for basin-level emissions are <5% of the estimated emission flux, and the DOFS are between 5 and 30 for the monthly inversion, indicating that the TROPOMI data are able to constrain basin-level methane emissions and partially resolve the spatial distribution on a monthly basis. The monthly variations in the posterior error and DOFS are mainly driven by uneven data coverage from TROPOMI sampling. For example, poor data coverage in November 2018 results in a large posterior error and a small DOFS (fig. S5).

We also perform an ensemble of sensitivity inversions by perturbing the configurations and parameters in the base inversion (table S2), aiming to characterize the uncertainties resulting from assumptions made in the inversion not captured by the analytical posterior error. Our results show that all these sensitivity inversions lead to consistent basin-level emission estimates. Annual mean fluxes from sensitivity inversions are within 0.5 Tg a⁻¹ of that from our base inversion (table S2), with general agreement in monthly variations as well (fig. S5). Because the uncertainty resulting from sensitivity inversions are significantly larger than that deduced from posterior error covari- \equiv the matrix (fig. S5), we report the uncertainty of our basin-ley mission estimate (0.5 Tg a⁻¹) as half of the range π om the inversion ensemble (2.4 to 3.4 Tg a⁻¹).

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Furthermore, to assess the uncertainty due to model transport o against measurements at the Midland Airport (MAF) in the Pe March 2019. Airport wind measurements are not assimilated i 🖉 GEOS-FP reanalysis (<u>45</u>), so these observations are independent. We find that the GEOS-FP 10-m wind speed both daytime and nighttime (fig. S8), with mean biases of less Ш errors in the model wind fields are unlikely to be a major source

compare hourly GEOS-FP 10-m wind speed n Basin during the period of May 2018 and bares well with the airport measurements in 6% in the mean wind speed. We conclude that error in the inversion.

We introduced a regional model bias term in monthly inversions to correct for regional background biases in simulated methane concentrations, which result mainly from imperfect boundary conditions. To check our estimate for this regional bias term, we sample the model simulation to compare with independent observations, i.e., surface measurements at the Mauna Loa Observatory (MLO; a Pacific free tropospheric site upwind of the North American continent) (<u>46</u>), tower measurements at Moody, Texas (WKT) (<u>47</u>), and aircraft measurements offshore Corpus Christi, Texas (TGC) (<u>48</u>). The latter two sites are geographically much closer to the Permian Basin (~400 km from WKT and ~700 km from TGC) than MLO, but can be affected by local emissions that are not optimized in our inversion. Our results show that the model simulation, when corrected with monthly regional model biases (derived from monthly inversions over the Permian Basin), is able to capture the observed monthly variation in methane concentrations, notably the sharp increase from August to October 2018 in MLO and WKT observations (fig. S9), supporting that it is necessary to optimize the regional model bias in the inversion. Better agreement is observed at MLO and TGC compared to WKT (fig. S9), likely because WKT is located closer to local sources that are not fully

optimized in the inversion. Overall, most of the differences between the prior simulation and TROPOMI observations can be explained by the regional model biases, except for the mismatch in the vicinity of the Permian Basin (fig. S2). We further perform a sensitivity inversion with a varied spatial domain (Bg Large). Compared to the base inversion, Bg Large results in a lower regional methane background (by 3 ppby on average) and a higher methane emission flux (3.4 Tg a^{-1}) (table S2 and fig. S5), reflecting the error correlation between regional methane biases and methane emissions.

In addition, we note that the inversion cannot fully explain the methane enhancement extending outside the Delaware Basin in the northwest direction (near 33°N, 105°W), although the inversion overall substantially improves the agreement between observations and model simulations (fig. S2). While our investigations do not attribute an obvious source of emissions causing the northwestern enhancement (whether oil/gas or other sources), the basin-level O/G emission estimates presented here are robust if this enhancement is caused by non-O/G sources, but are conservative if it is caused by O/G sources.

Emission inventory based on bottom-up information

We create a bottom-up methane emission estimate (EI_{BU}) for the study domain starting from the gridded version of the EPA anthropogenic greenhouse gas emission inventory for 2012 (<u>49</u>). Maasakkers *et al.* (<u>49</u>) developed a procedure to spatially and temporally allocate the national sectorial methane emissions reported in the U.S. Inventory of Greenhouse Gas Emissions and Sinks (GHGI) by U.S. EPA on a $0.1^{\circ} \times 0.1^{\circ}$ grid, using various databases at the state, county, local, and point-source level. The emission inventory includes methane emissions from agriculture, coal mining, natural gas systems, petroleum (oil) systems, waste, and other minor anthropogenic sources.

To reflect the intensifying exploitation activity in recent years in the Permian Basin, we then make an extrapolation of the methane emissions from the oil and gas production sector, using 2018 Enverus Drillinginfo data on well count, well completion, and production (50). To account for the changes in the national average emission factors, we further scale the subsectorial production emissions using the ratio between the latest GHGI (22) and a previous GHGI that Maasakkers *et al.* (49) was based on (51) for 2013 emissions. The updates result in total methane emissions of 1.2 Tg a⁻¹ in the Permian Basin (blue box in Fig. 5A), with 1.0 Tg a⁻¹ coming from O/G-related emissions and the remainder mainly from agriculture. We use this updated gridded emission inventory (EI_{BU}) as the prior emission estimate for the inversion. The resulting emissions inventory dataset (EI_{BU} inventory) is publicly available for our study region encompassing the entire Permian Basin (<u>https://doi.org/10.7910/DVN/NWQGHU</u>).

Emission inventory based on site-level emission measurements

An alternative prior estimation of methane emissions is obtained by extrapolating ground-based methane emis-⊨ on measurements from a limited sample of oil and gas produt n sites in the Permian Basin (primarily in the . The measurements found a wide range of site-New Mexico portion of the basin) during July and August 2018 level emission rates, which appear to be associated with the co emission rates for simple (with only wellheads and/or pump ja • versus complex sites (also with storage tanks and/or compressors). Extrapolating these site-level emission r ane emission rate of 2.3 Tg a⁻¹ from O/G production. Addition issions from compressor stations and processing plants are estimated to be 0.22 and 0.14 Tg a⁻¹, respectivel 📃 ing activity data from Enverus Drillinginfo's midstream infrastructure dataset, facility-level emission facto in literature (<u>53</u>, <u>54</u>), and blowdown event emission factors from GHGI (<u>22</u>). We then disaggregate the ba evel O/G-related emissions to a $0.1^{\circ} \times 0.1^{\circ}$ grid by the spatial distribution of gas production (Fig. 2D). To comp 🗲 the inventory, non-O/G anthropogenic methane emissions (0.2 Tg a^{-1}) are taken from EI_{BU}. This emission inventory (EI_{ME}), based primarily on extrapolation of limited site-level measurements, provides an alternative prior estimate for the inversion and is used to test the sensitivity of the results to the choice of prior information (fig. S3). See text S2 for detailed information regarding the site-level measurements and the extrapolation procedure. The resulting emissions inventory dataset (EI_{ME} inventory) is publicly available for our study region encompassing the entire Permian Basin (https://doi.org/10.7910/DVN/NWQGHU).

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Supplementary Material

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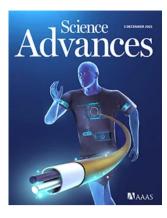
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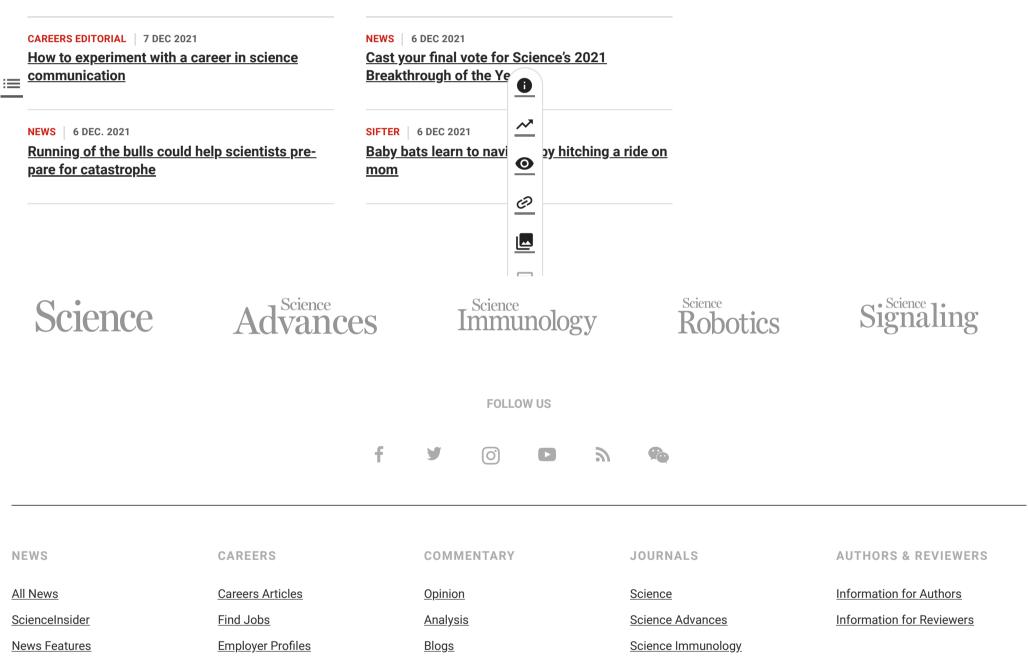
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